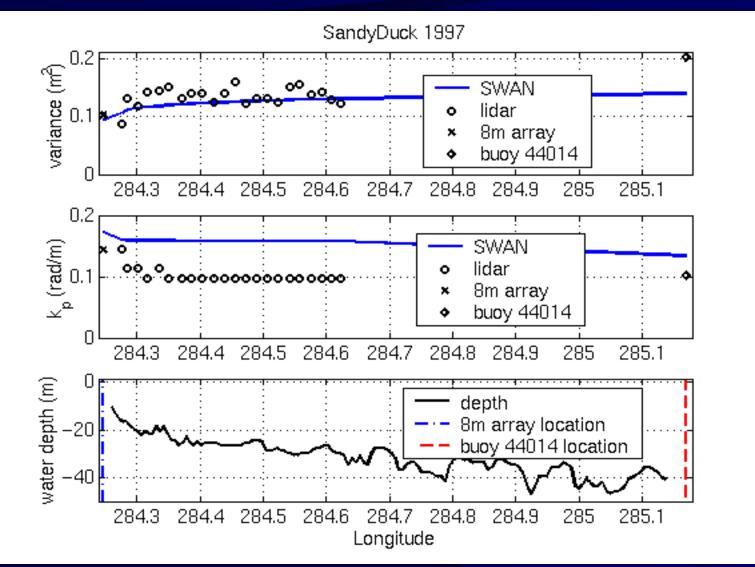
# SWAN model analysis using a Lake Michigan storm event

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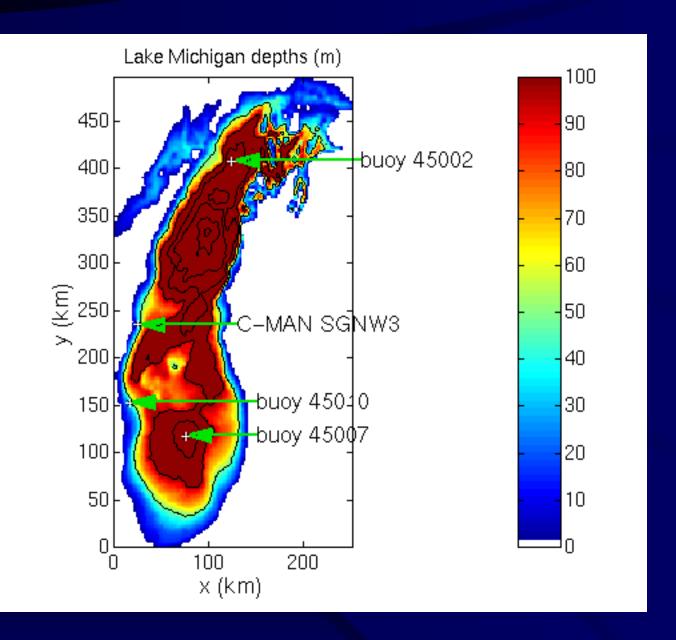
#### Motivation



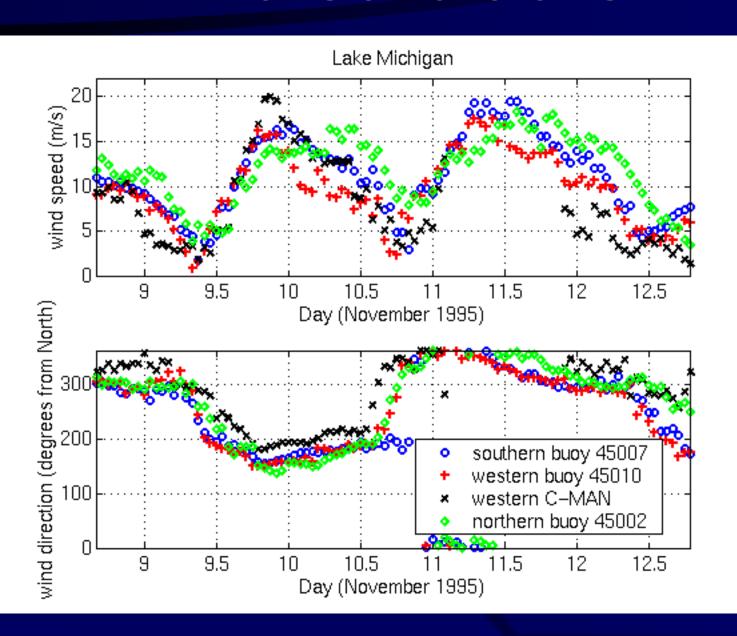
Concern: open ocean grid boundary (~400 km offsh

#### Lake Michigan, Nov. 1995

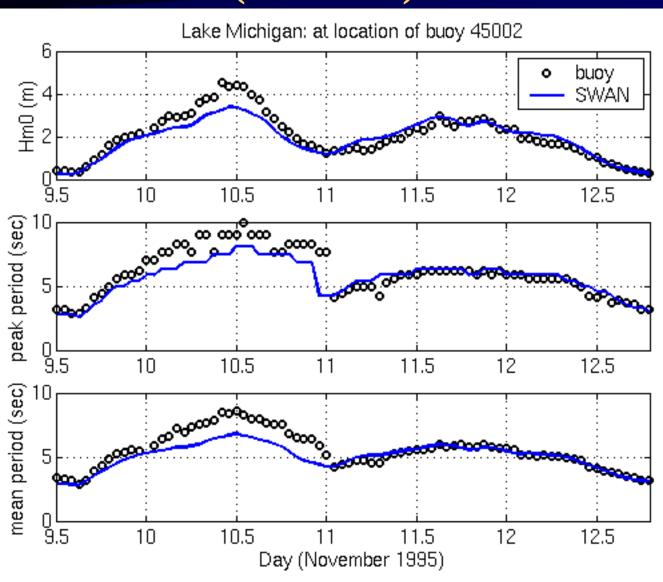
- Two-part storm event
- Forcing: wind data from 2 deep water NDBC buoys
- Computational resolution: 2km
- Output comparisons: deep water only



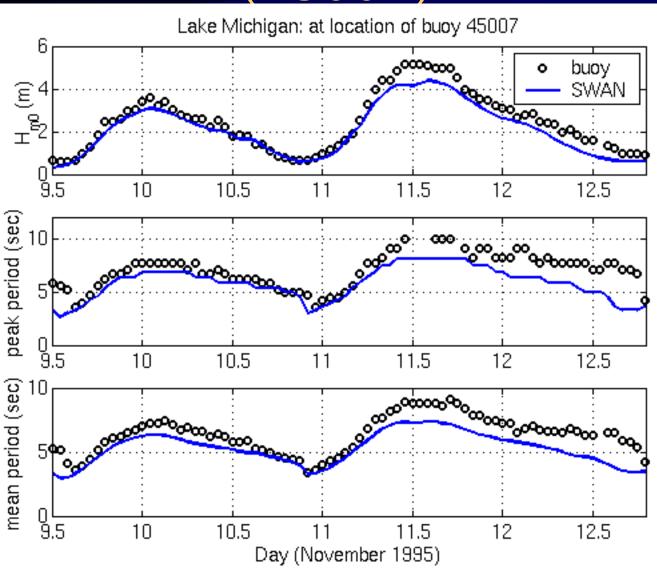
#### Wind Conditions



### Results at northern buoy (45002)



### Results at southern buoy (45007)



### Alteration of deepwater source/sink terms

- Dissipation term is least well understood; arguably the least accurate. ⇒ greatest potential payoff
- work within framework of Komen et al. (1984) formulation, or create something completely new?
  - We take the former approach

#### Komen et al. (1984) dissipation mechanism:

 $\beta_{Diss}E(\sigma,\theta)$ : dissipation of wave action

$$eta_{Diss}(\sigma, heta) = C_{ds} \left[ \frac{s}{s} \right]^{m} \sigma_{m} \left[ \frac{k}{k_{m}} \right]^{n}$$

n affects the weighting of dissipation toward higher or frequencies. n=1 is traditionally used.

This choice is based on a) tuning to data in *fully-developments* and b) theoretical arguments of uncertain when we find that optimally, 1 < n < 1.5.

We observe the expected effect of n on spectral shape:

A modified n requires a modified  $C_{ds}$  (for agreement with Thus, the effect of n on total wave energy tends to be s

# Komen et al. (1984) dissipation mechanism: Issue #2

- The formulation is meant to represent the process of steepness-limited breaking.
- Under normal circumstances, swell does not break.
- We implement a switch to prevent this.
- Significant improvement in model skill.
- Care must be taken to allow for situations in which swell *is* expected to break.

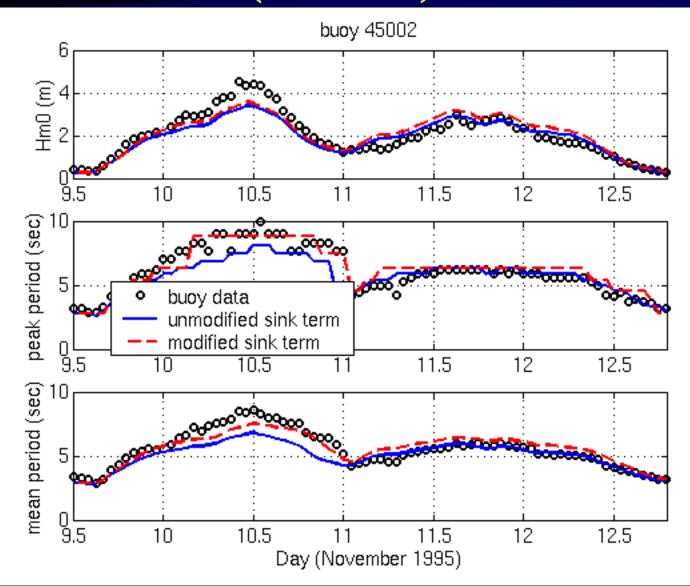
## Komen et al. (1984) dissipation mechanism: Issue #3

- Dissipation at a given frequency is dependent on the mean steepness of the entire wave spectrum
  - Can produce aphysical results (e.g. as demonstrated by van Vledder (1999))
  - Might be addressed by treating dissipation of wave components separately.

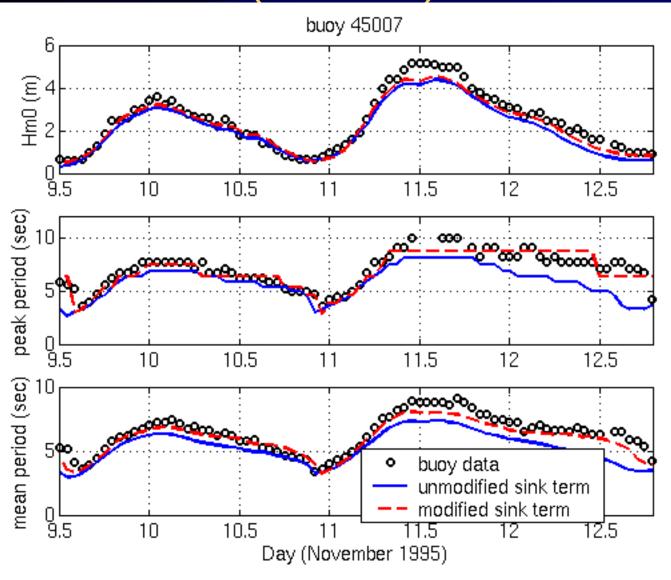
#### Calculated relative error values for Lake Michigan simulations

		NDBC buoy 45002			NDBC buoy 45007		
Param eter	Swell breaks?	n=1	n=1.5	n=2	n=1	n=1.5	n=2
$H_{m0}$	yes	0.18	0.22	0.25	0.16	0.22	0.26
	no	0.23	0.18	0.19	0.09	0.11	0.18
$T_{mean}$	yes	0.15	0.11	0.09	0.19	0.13	0.09
	no	0.12	0.10	0.09	0.12	0.09	0.07
T <sub>peak</sub>	yes	0.15	0.12	0.12	0.22	0.16	0.11
	no	0.14	0.11	0.12	0.11	0.10	0.10

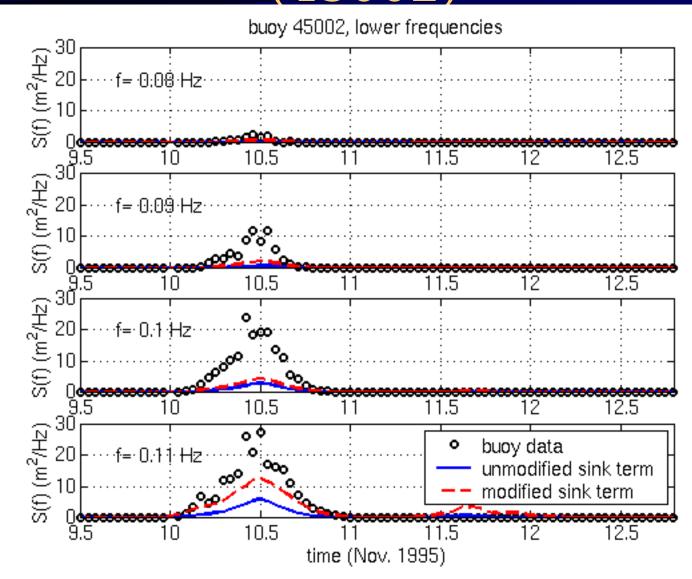
### New results at northern buoy (45002)



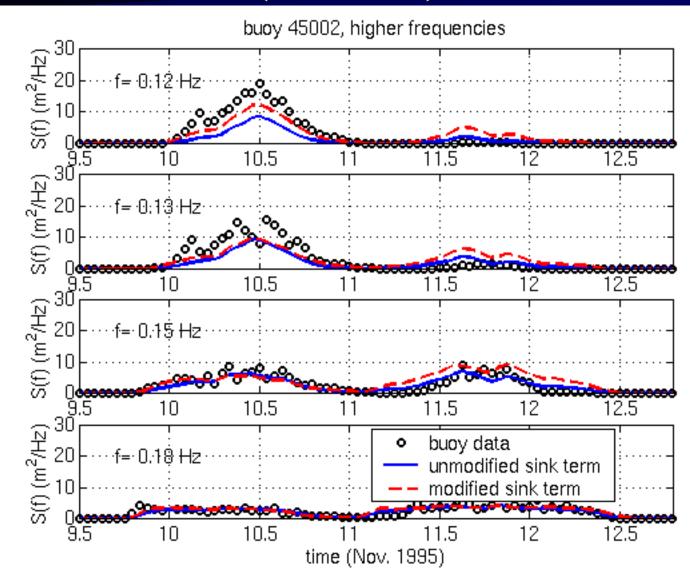
### New results at southern buoy (45007)



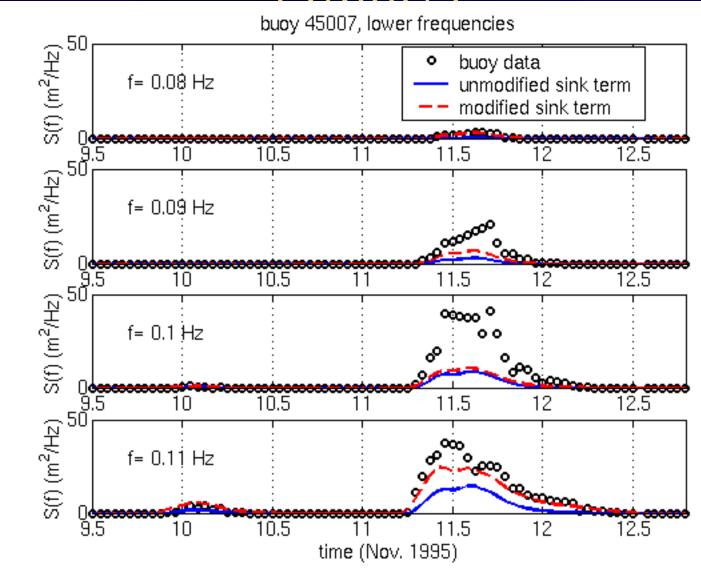
### New results at northern buoy (45002)



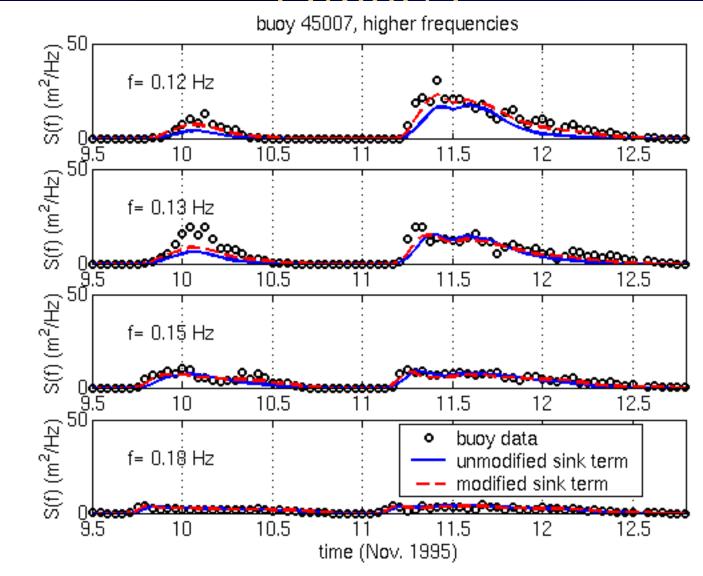
### New results at northern buoy (45002)

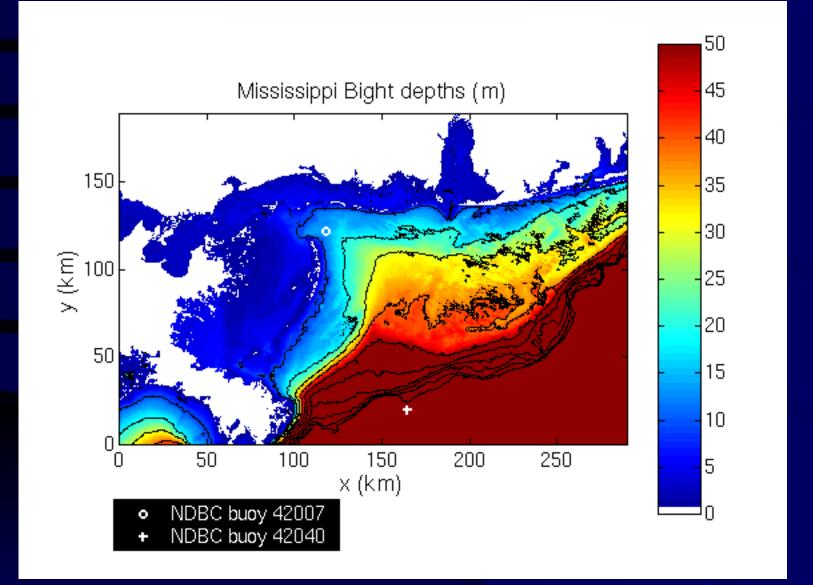


#### New results at southern buoy (45007)

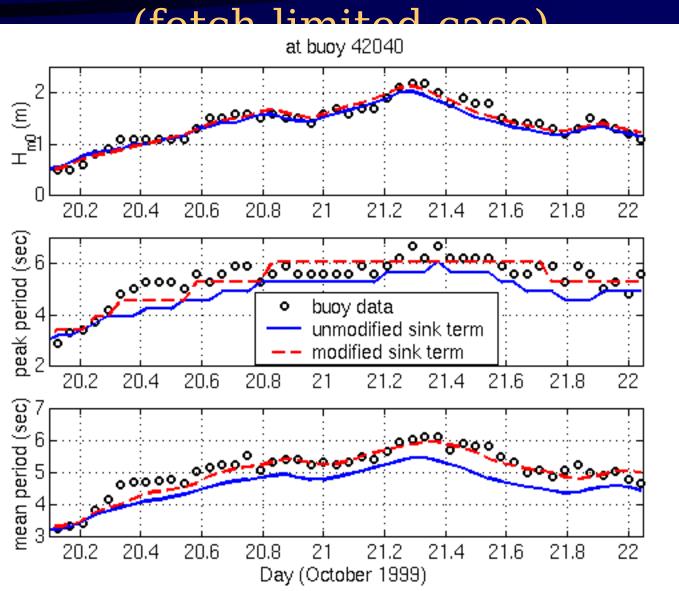


#### New results at southern buoy (45007)

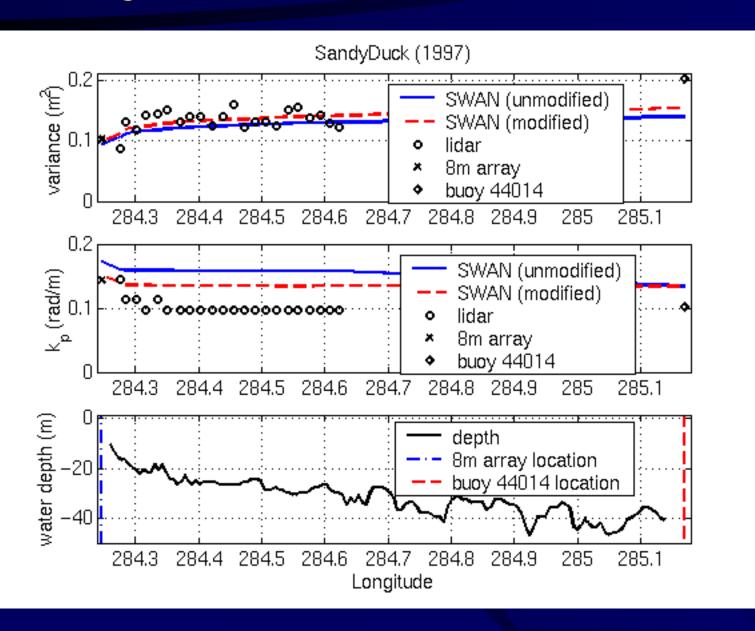




#### MS Bight: Comparison between NDBC buoy 42040 and SWAN



#### SandyDuck (1997): new results



#### Summary

- alteration of *n* (weighting of dissipation toward higher or lower frequencies): minor effect on results (due to P-M "shackles").
- "swell does not break under normal conditions" rule: significant pay-off.
- Aphysical dependence of wave breaking on total steepness: not addressed yet.

#### Concerns

- The other two source terms (e.g. DIA)
- Tail representation (e.g. Banner and Young (1994))
- Wind forcing
- $S_{wcap}$ : may be better to start over from scratch.
- Directional behavior